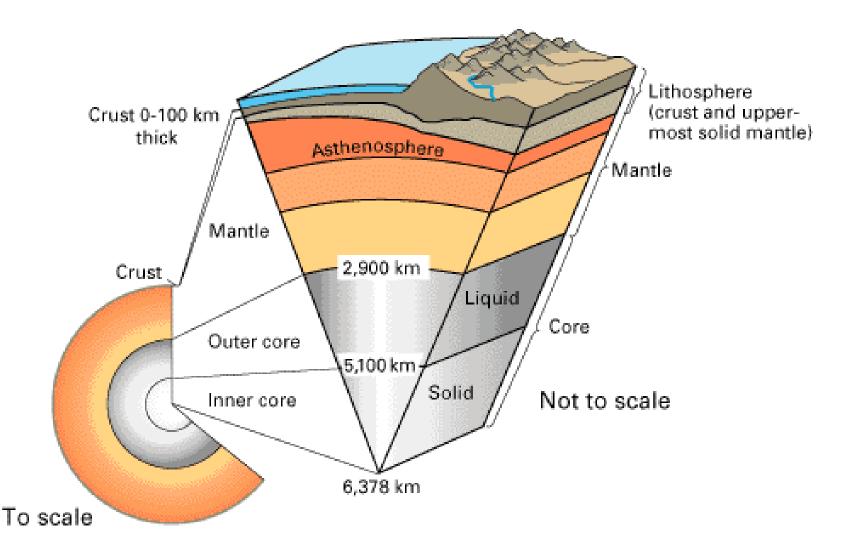
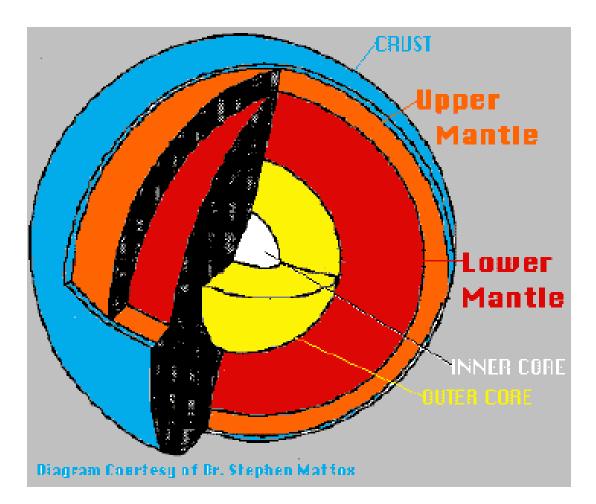
The Layers of the Earth

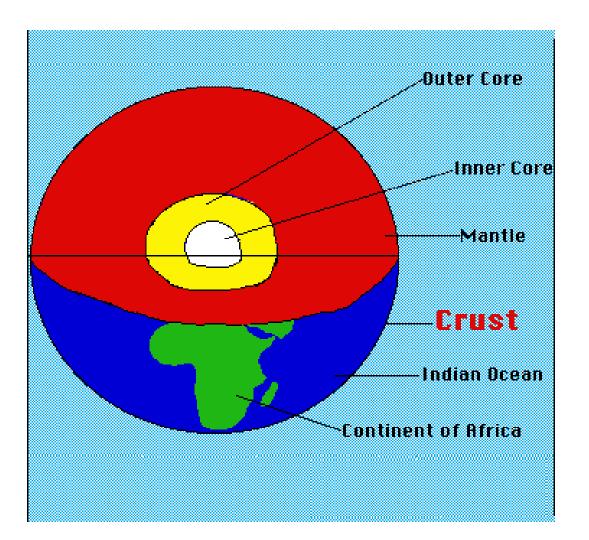


The Four Layers



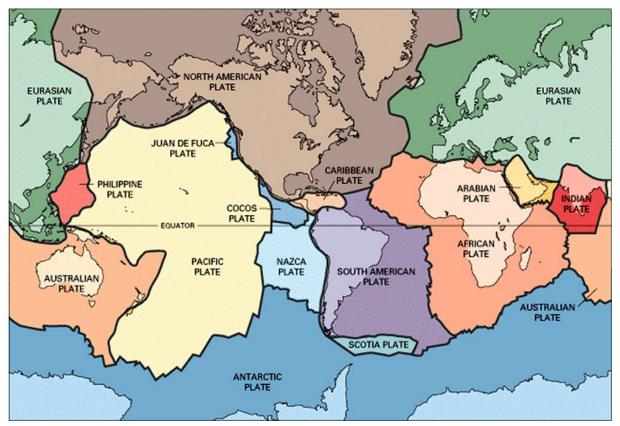
The Earth is composed of four different layers. The **crust** is the layer that you live on, and it is the most widely studied and understood. The **mantle** is much hotter and has the ability to flow. The **outer** core and inner core are even hotter with pressures so great you would be squeezed into a ball smaller than a marble if you were able to go to the center of the Earth!

The Crust



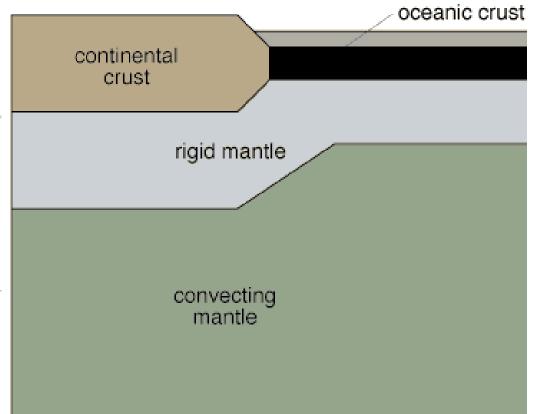
The Earth's **Crust** is like the skin of an apple. It is very thin in comparison to the other three layers. The crust is only about 3-5 miles (8 kilometers) thick under the oceans (**oceanic crust**) and about 25 miles (32 kilometers) thick under the continents (**continental crust**).

The Lithospheric Plates



The **crust** of the Earth is broken into many pieces called **plates**. The plates "float" on the soft, semi-rigid **asthenosphere**.

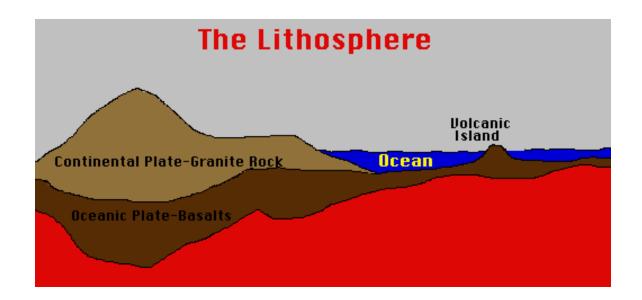
The Asthenosphere



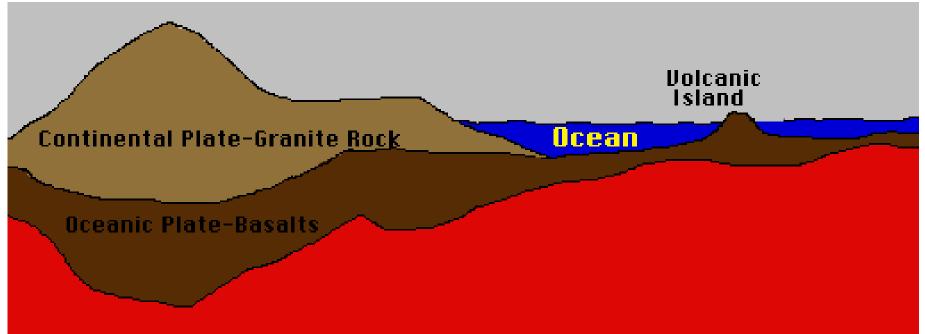
The **asthenosphere** is the semi-rigid part of the **middle mantle** that flows like hot asphalt under a heavy weight.

The Lithosphere

The **crust and the upper layer of the mantle** together make up a zone of rigid, brittle rock called the **Lithosphere**.

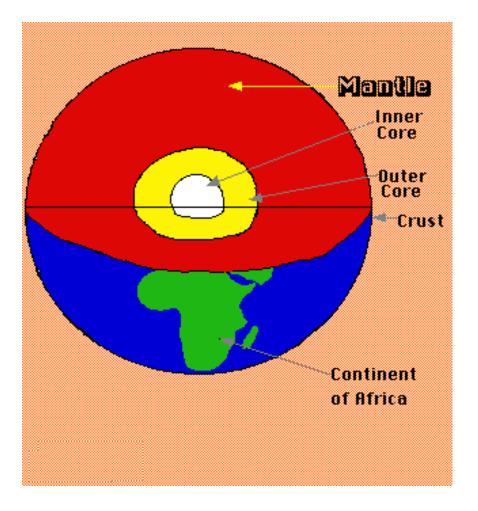


The Crust



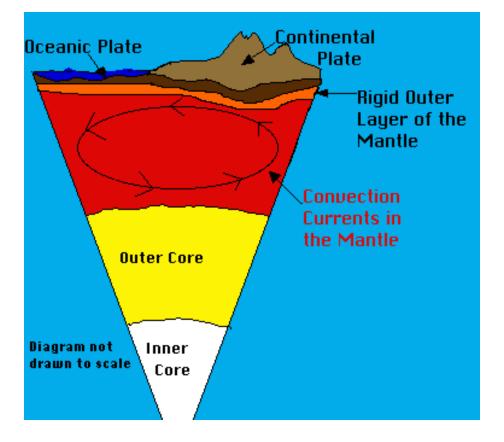
The **crust** is composed of two rocks. The **continental crust** is mostly **granite**. The **oceanic crust** is **basalt**. Basalt is much denser than the granite. Because of this the less dense continents ride on the denser oceanic plates.

The Mantle



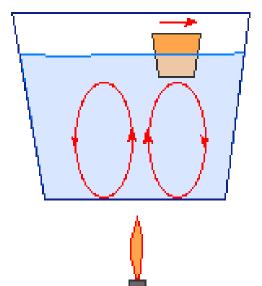
The **Mantle** is the largest layer of the Earth. The **middle mantle** is composed of very hot dense rock that flows like asphalt under a heavy weight. The movement of the middle mantle (asthenosphere) is the reason that the crustal plates of the Earth move.

Convection Currents



The middle mantle "flows" because of convection currents. **Convection** currents are caused by the very hot material at the deepest part of the mantle rising, then cooling and sinking again --repeating this cycle over and over.

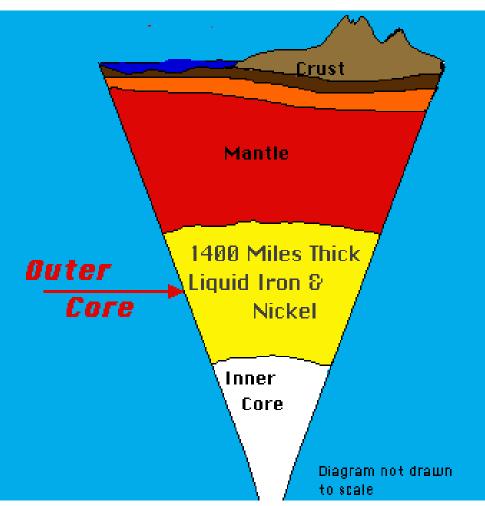
Convection Currents



The next time you heat anything like soup or water in a pan you can watch the **convection currents** move in the liquid. When the convection currents flow in the **asthenosphere** they also move the crust. The crust gets a free ride with these currents, like the **cork** in this illustration.

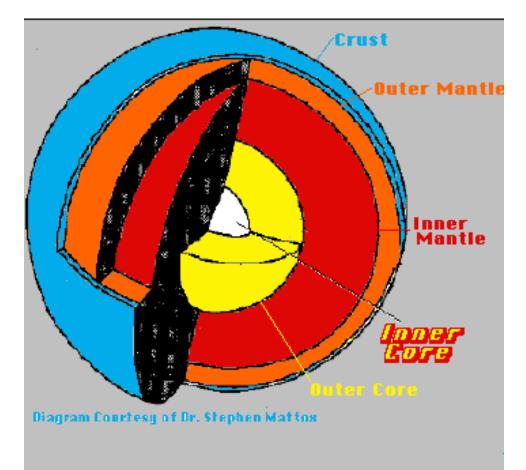
Safety Caution: Don't get your face too close to the boiling water!

The Outer Core



The core of the Earth is like a ball of very hot metals. The outer core is so hot that the metals in it are all in the liquid state. The outer core is composed of the melted metals of nickel and iron.

The Inner Core



The **inner core** of the Earth has temperatures and pressures so great that the metals are squeezed together and are not able to move about like a liquid, but are forced to vibrate in place like a **solid**.

Faults and Earthquakes

Take-Away Points

- 1. Earthquakes generate waves that travel through the earth
- 2. Earthquakes occur when rocks slip along faults
- 3. Faults are classified by the kinds of movement that occur along them
- 4. Earthquakes don't kill people, buildings kill people
- 5. Magnitude and Intensity
- 6. Seismic waves are used to map the earth's interior
- 7. Predicting earthquakes is not yet possible

Some Important Earthquakes

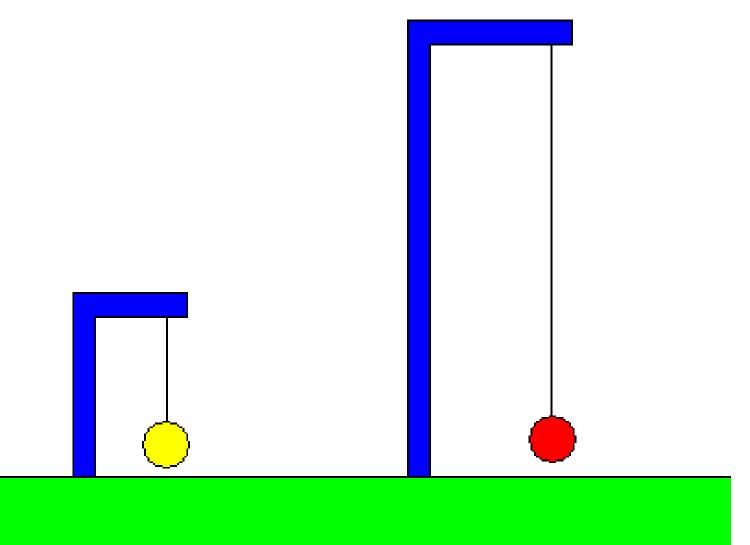
- 1755 Lisbon, Portugal
- Killed 70,000, Raised Waves in Lakes all over Europe
- First Scientifically Studied Earthquake 1811-1812 - New Madrid, Missouri
- Felt over 2/3 of the U.S.
- Few Casualties
- 1886 Charleston, South Carolina
- Felt All over East Coast, Killed Several Hundred.
- First Widely-known U.S. Earthquake

Some Important Earthquakes

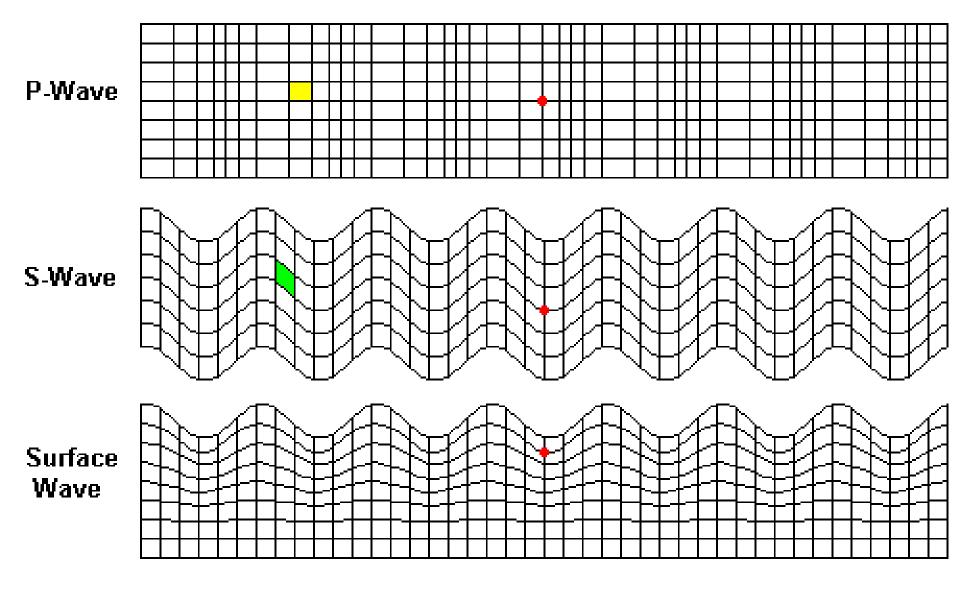
1906 - San Francisco

- Killed 500 (later studies, possibly 2,500)
- First Revealed Importance of Faults
 1923 Tokyo Killed 140,000 in firestorm
 1964 Alaska
- Killed about 200
- Wrecked Anchorage.
- Tsunamis on West Coast.
- 1976 Tangshan, China
- Hit an Urban Area of Ten Million People
- Killed 650,000

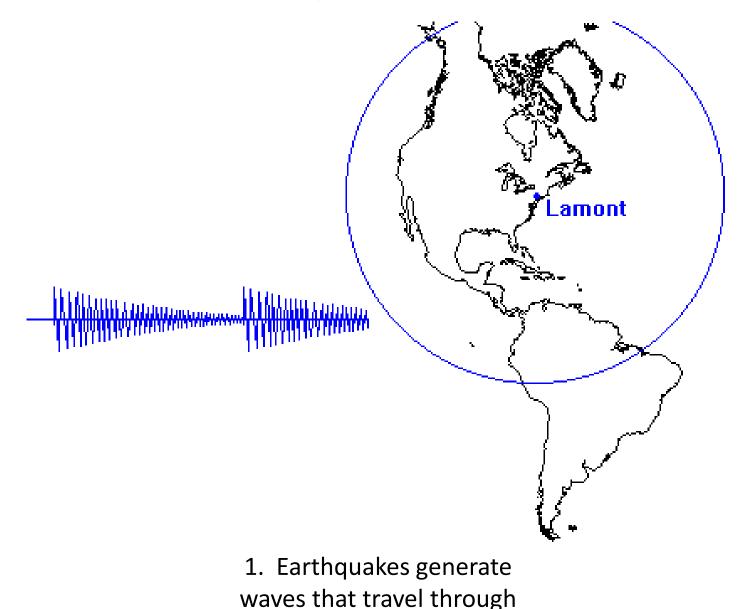
How Seismographs Work



Seismic Waves

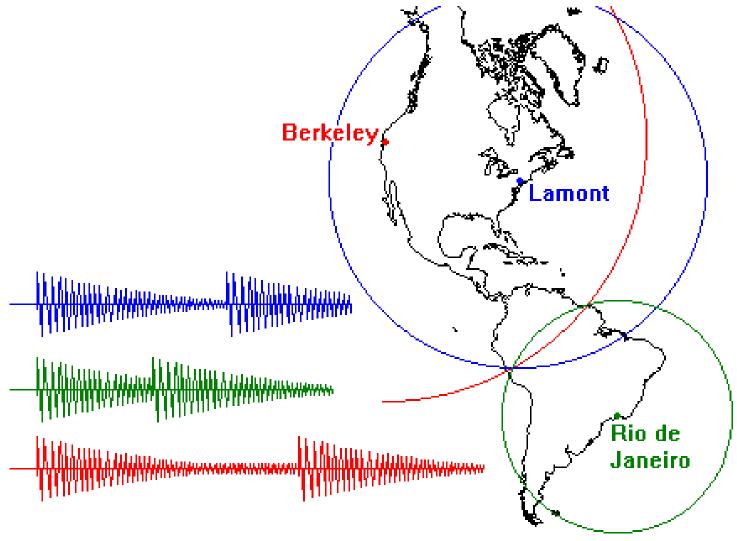


Locating Earthquakes

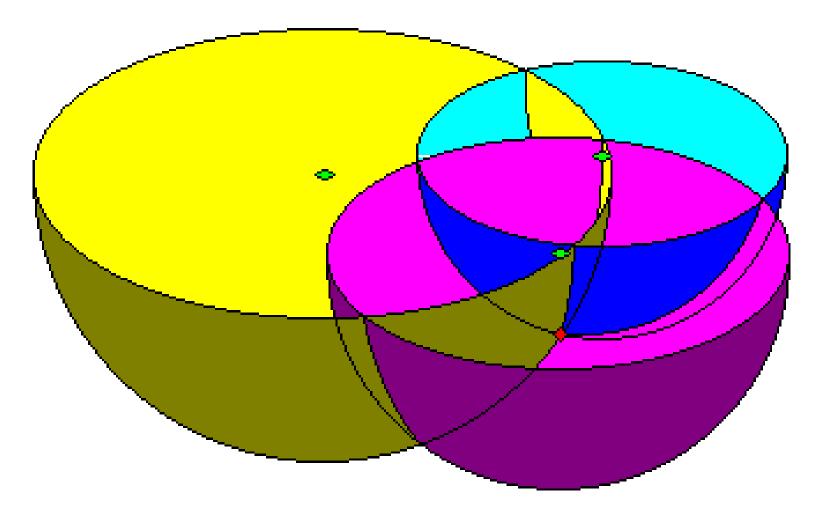


Locating Earthquakes ₽ _amont Rio de Janeiro 1. Earthquakes generate

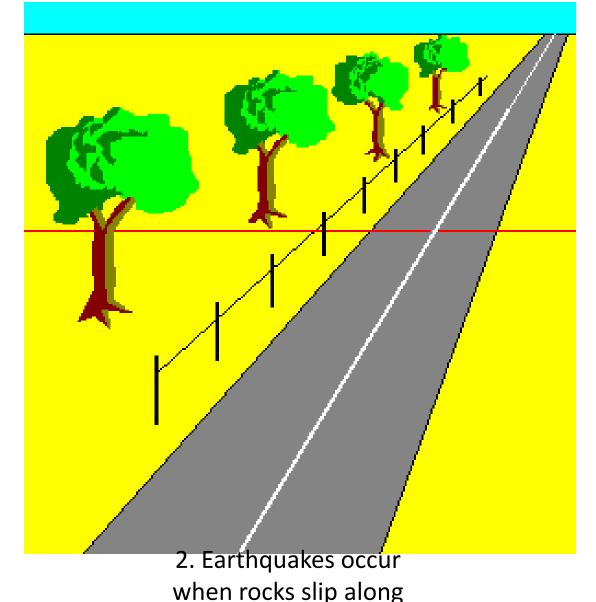
Locating Earthquakes



Locating Earthquakes - Depth



Elastic Rebound



Epicenter and Focus

Focus

• Location within the earth where fault rupture actually occurs

Epicenter

• Location on the surface above the focus

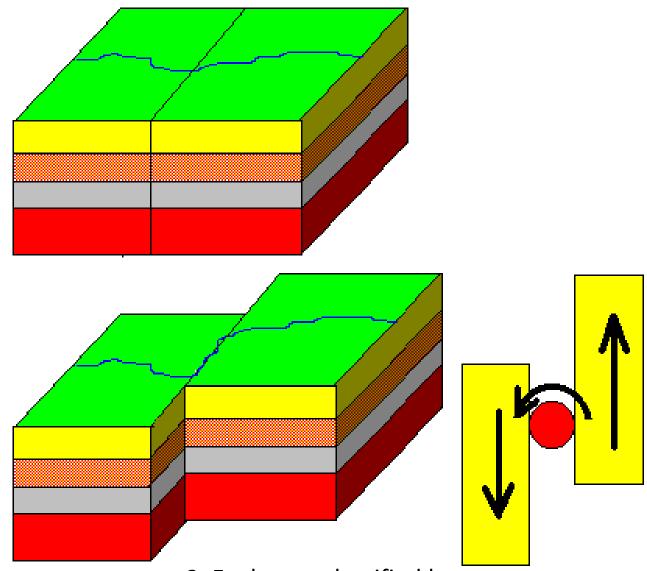
2. Earthquakes occur when rocks slip along

Types of Faults

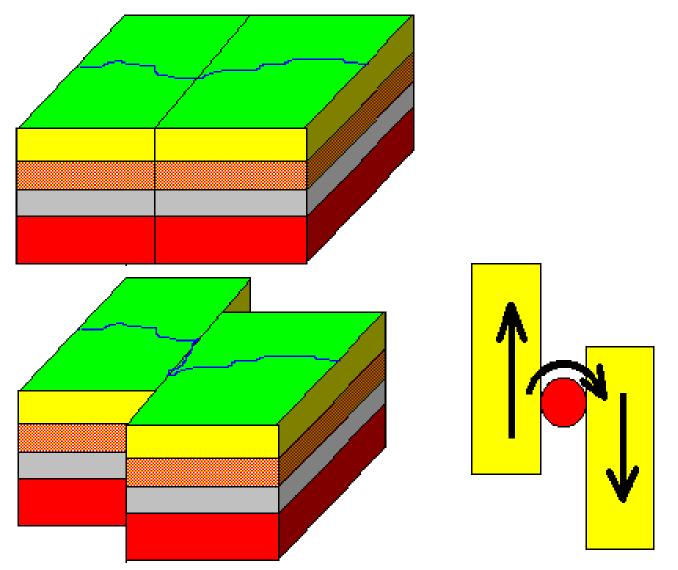
Faults Are Classified According to the Kind of Motion That Occurs on Them

- Joints No Movement
- Strike-Slip Horizontal Motion
- Dip-Slip Vertical Motion

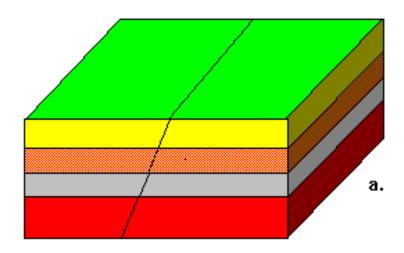
Strike-Slip Fault – Left Lateral



Strike-Slip Fault – Right Lateral

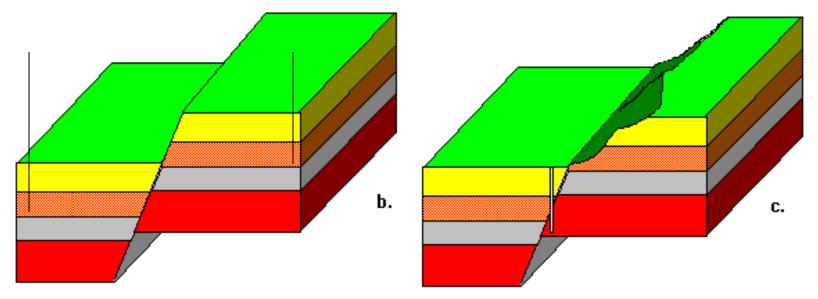


Dip-Slip Fault - Normal

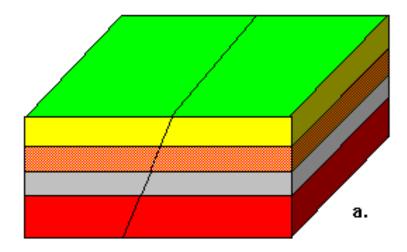


Normal Faulting

- a. A block of crust before faulting
- b. After faulting. Note that the block becomes longer.
- c. An eroded normal fault. Note that the well misses the gray layer completely.

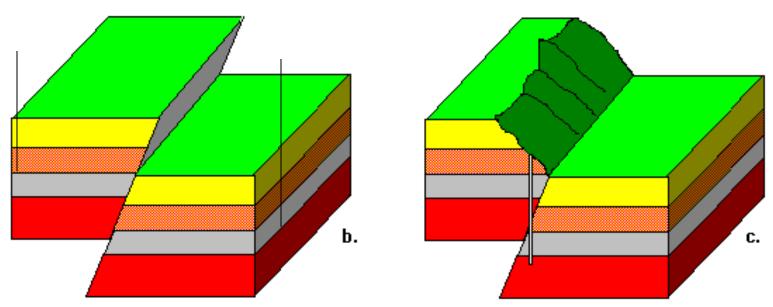


Dip-Slip Fault - Reverse



Reverse or Thrust Faulting

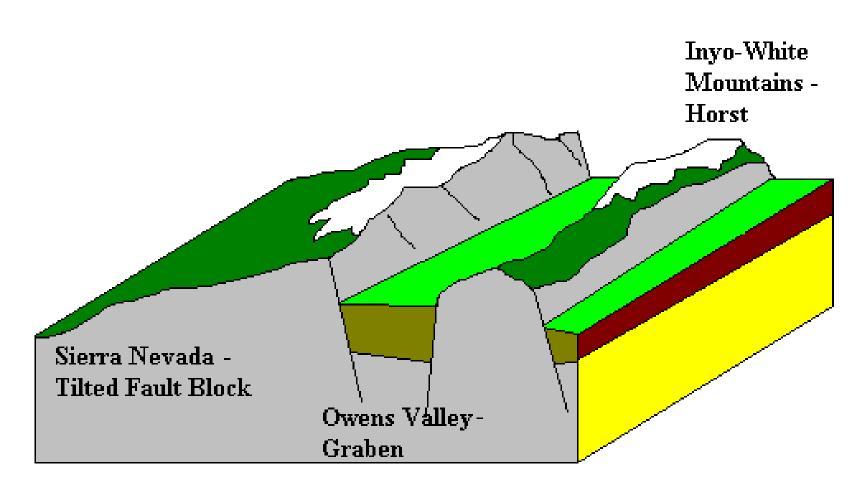
- a. A block of crust before faulting.
- b. After faulting. Note that the block becomes shorter.
- c. An eroded reverse fault. Note that the well passes through several layers twice.



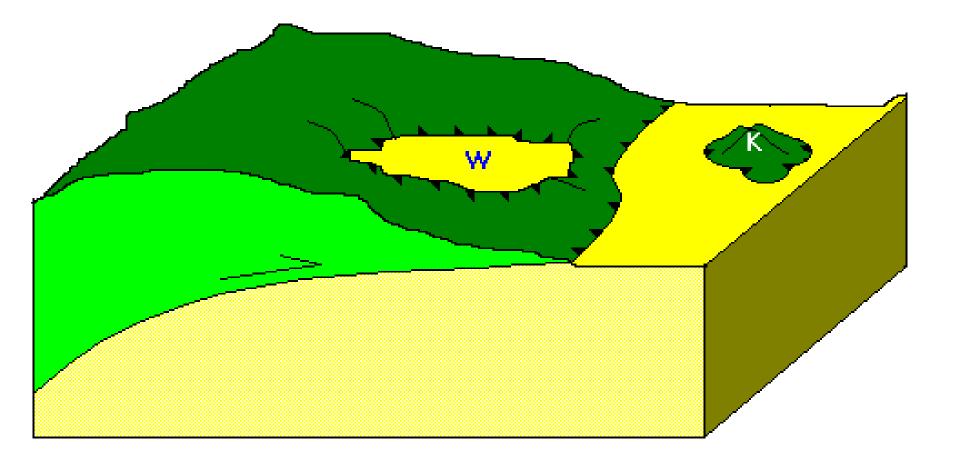
Dip-Slip Faults

- Normal Faults: Extension
- Reverse Faults: Compression
 - -Reverse Faults are often called Thrust Faults

Normal Fault Structures



Reverse Fault Structures



Major Hazards of Earthquakes

- Building Collapse
- Landslides
- Fire
- Tsunamis (Not Tidal Waves!)

4. Earthquakes don't kill people, buildings kill

Safest & Most Dangerous Buildings

- Small, Wood-frame House Safest
- Steel-Frame
- Reinforced Concrete
- Unreinforced Masonry
- Adobe Most Dangerous

4. Earthquakes don't kill people, buildings kill

Tsunamis

Probably Caused by Submarine Landslides Travel about 400 M.p.h.

Pass Unnoticed at Sea, Cause Damage on Shore

Warning Network Around Pacific Can Forecast Arrival

Whether or Not Damage Occurs Depends on:

- Direction of Travel
- Harbor Shape
- Bottom
- Tide & Weather

4. Earthquakes don't kill people, buildings kill

Magnitude and Intensity

Intensity

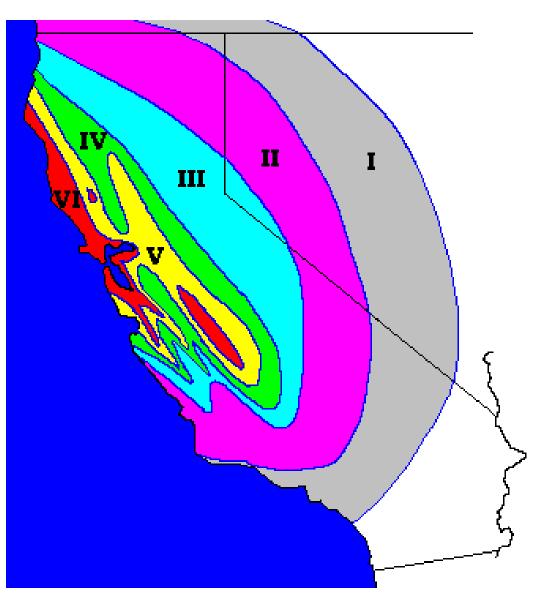
- How Strong Earthquake Feels to Observer Magnitude
- Related to Energy Release
- Determined from Seismic Records
- Rough correlation between the two for shallow earthquakes

Intensity

How Strong Earthquake Feels to Observer Depends On:

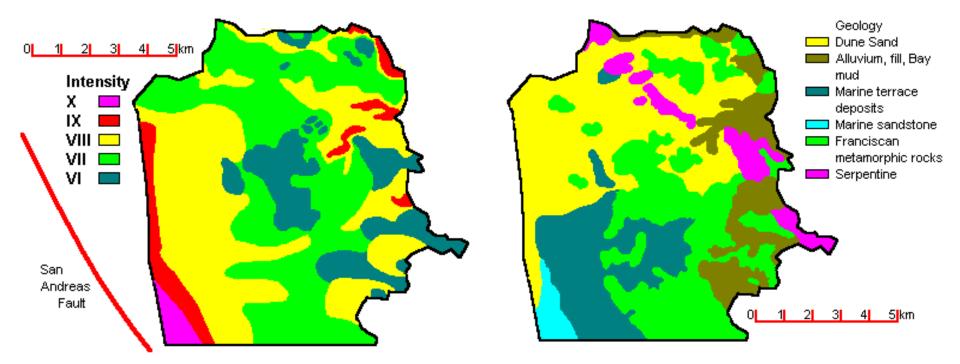
- Distance to Quake
- Geology
- Type of Building
- Observer!
- Varies from Place to Place
- Mercalli Scale- 1 to 12

5. Magnitude and Intensity Isoseismals from the 1906 San Francisco Earthquake



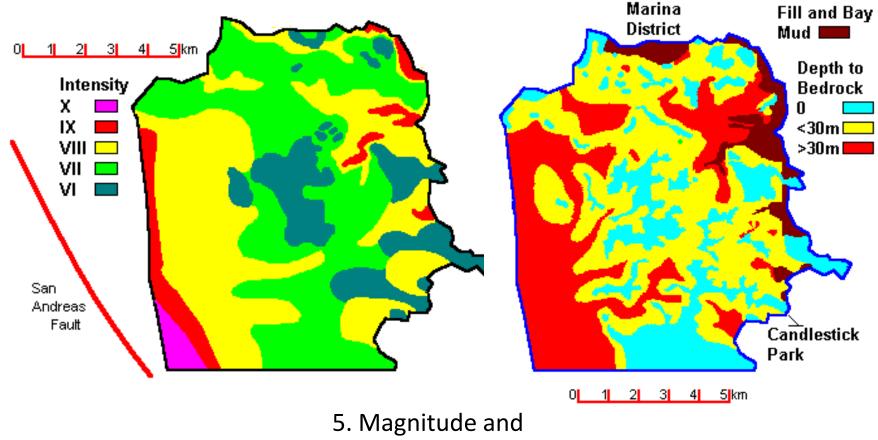
5. Magnitude and Intensity

Intensity and Geology in San Francisco, 1906



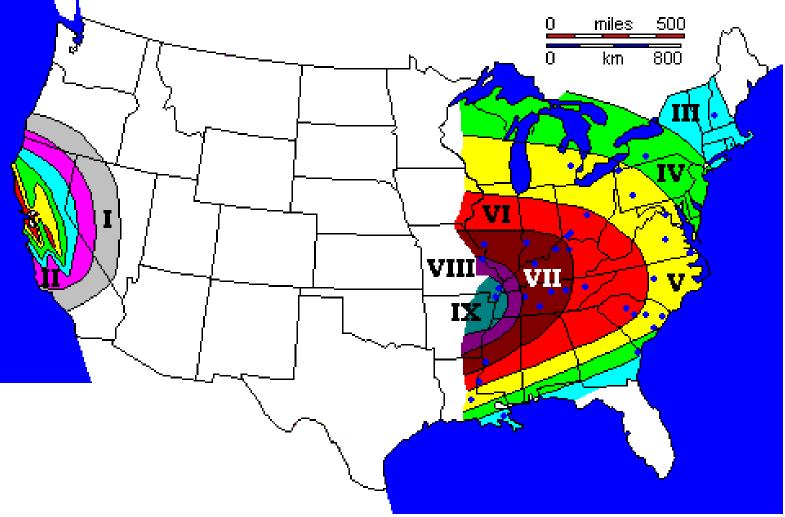
5. Magnitude and Intensity

Intensity and Bedrock Depth in San Francisco, 1906



Intensity

San Francisco and New Madrid Compared



5. Magnitude and Intensity

Magnitude - Determined from Seismic Records

Richter Scale:

- Related to Energy Release
- Exponential
- No Upper or Lower Bounds
- Largest Quakes about Mag. 8.7

- Magnitude-Energy Relation
 - 4 1
 - 5 30
 - 6 900:
 - -1 Megaton = about 7
 - 7 27,000
 - 8 810,000

5. Magnitude and Intensity

Seismic - Moment Magnitude

- A Seismograph Measures Ground Motion at One Instant But --
- A Really Great Earthquake Lasts Minutes
- Releases Energy over Hundreds of Kilometers
- Need to Sum Energy of Entire Record
- Modifies Richter Scale, doesn't replace it
- Adds about 1 Mag. To 8+ Quakes

5. Magnitude and Intensity

Magnitude and Energy

Magnitude	Energy	Explosive Power	Example	
9	U.S. Energy Use for a month		Alaska 1964 Indonesia 2004	
8	U.S. Energy Use for a day		San Francisco, 1906	
7		One Megaton	World Series Earthquake, 1989	
6	U.S. Energy Use for a minute	Large Thunderstorm		
5		One Kiloton		
4				
3		One ton of explosives	World Trade Center Collapse	
5. Magnitude and				
Intensity				

Magnitude and Energy

	-			
Magnitude	Energy	Explosive Power	Example	
3		One ton of Explosives	World Trade Center Collapse	
2				
1	Topple 50-meter tree	One kilogram of explosives	Head-on colision at 60 mph	
0	Drop a car 10 meters	Half stick of dynamite	Very bad day skydiving	
-1	Impact of bullet	One gram of explosives		
-2	Hammer blow			
-3	Dribbling a basketball			
5. Magnitude and				

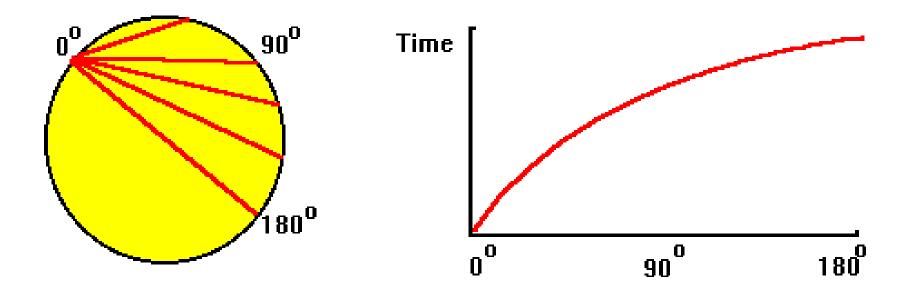
Intensity

Seismology and Earth's Interior

Successive Approximation in Action

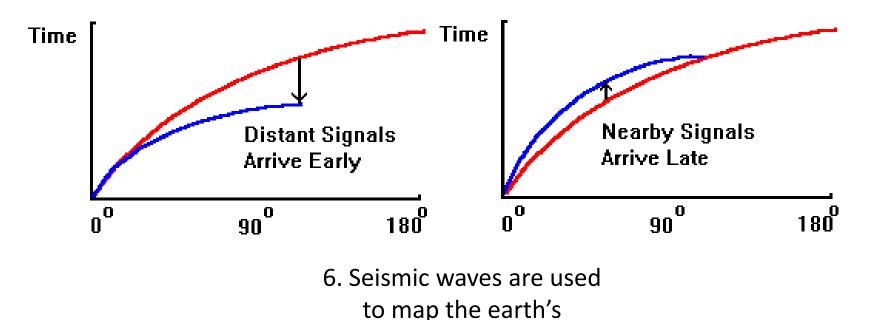
1. Assume the Earth is uniform.

• We know it isn't, but it's a useful place to start. It's a simple matter to predict when a seismic signal will travel any given distance.



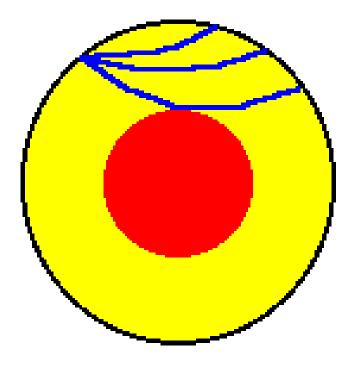
2. Actual seismic signals don't match the predictions

- If we match the arrival times of nearby signals, distant signals arrive too soon
- If we match the arrival times of distant signals, nearby signals arrive too late.
- Signals are interrupted beyond about 109 degrees

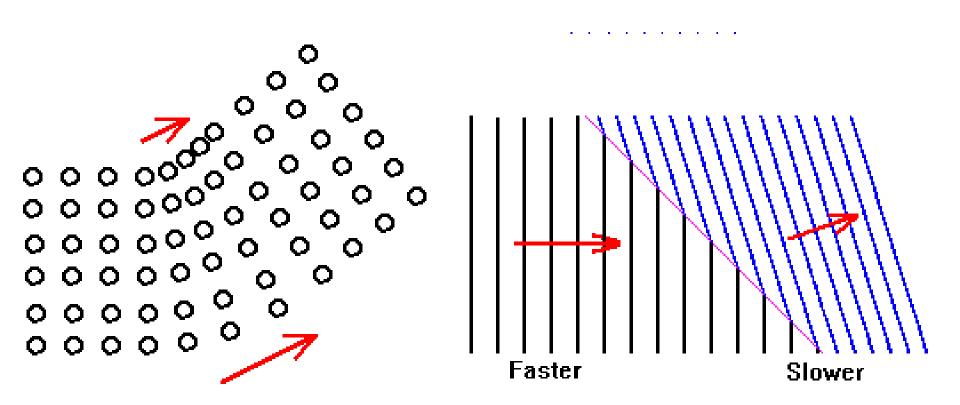


3. We conclude:

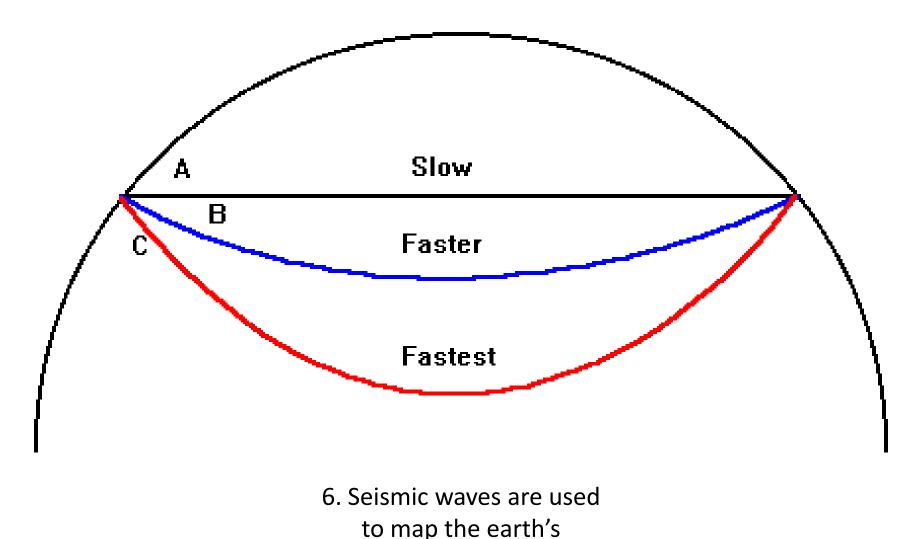
- Distant signals travel through deeper parts of the Earth, therefore ..
- 2. Seismic waves travel faster through deeper parts of the Earth, therefore
- 3. They travel curving paths (refract)
- 4. Also, there is an obstacle in the center (the core).



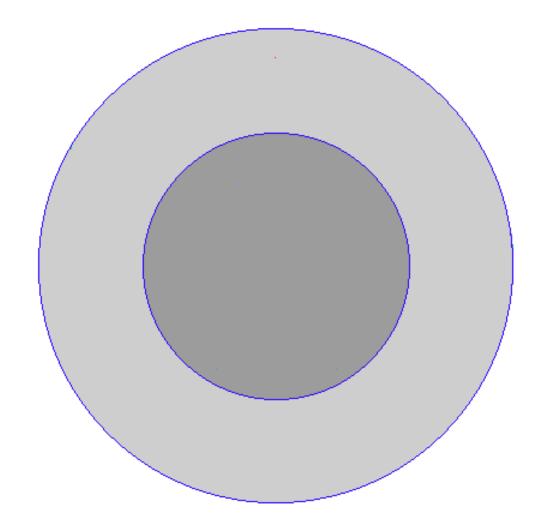
Why Refraction Occurs



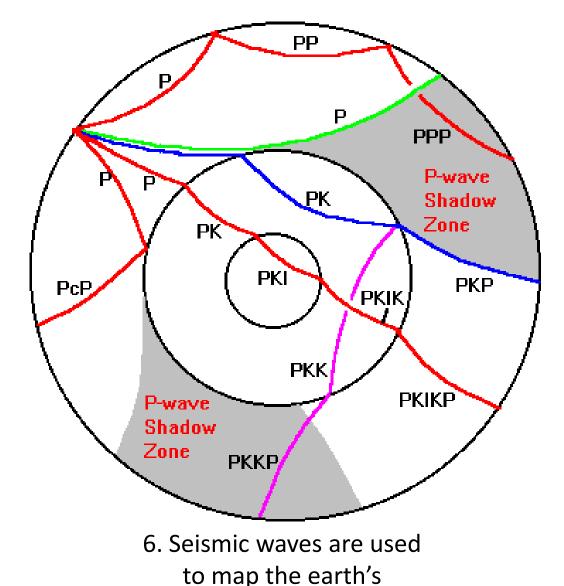
Waves Travel The Fastest Path



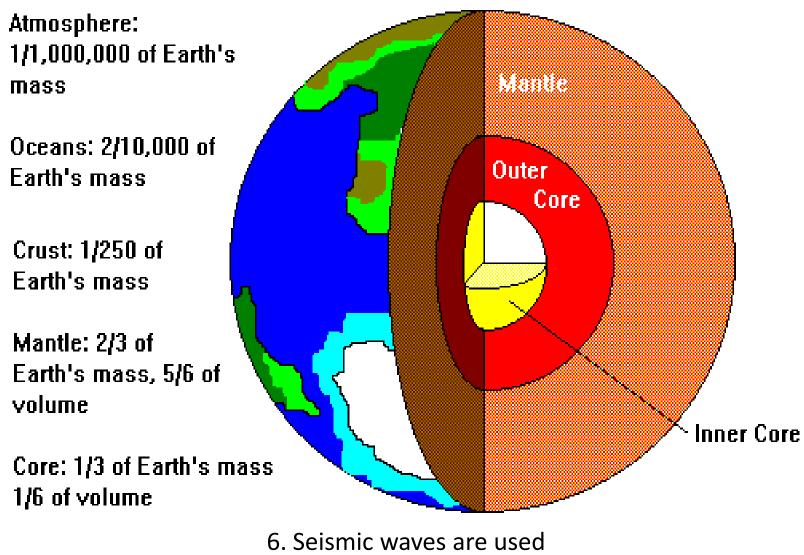
Seismic Waves in the Earth



Inner Structure of the Earth



The overall structure of the Earth



to map the earth's

Strategies of Earthquake Prediction

- Lengthen Historical Data Base
- Historical Records
- Paleoseismology
 Short-term Prediction
- Precursors

Long-term Prediction

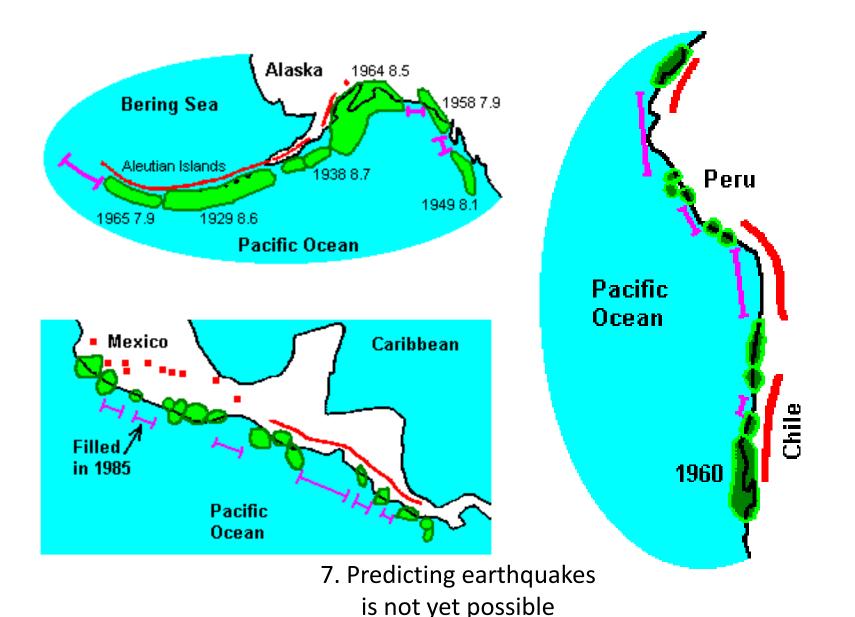
- Seismic Gaps
- Risk Levels

Modeling

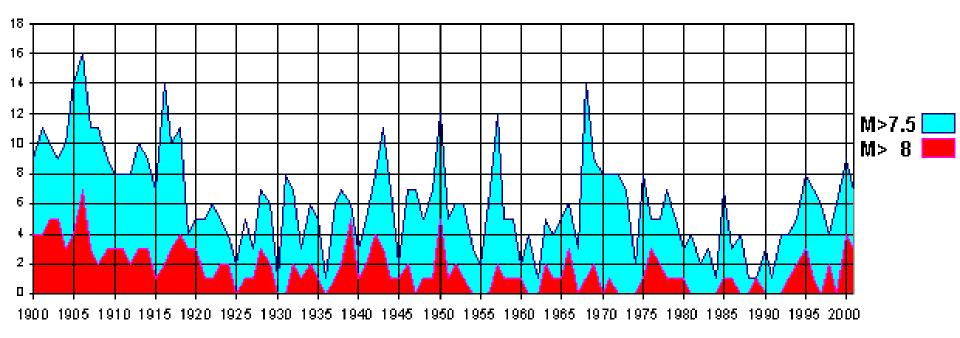
- Dilatancy Diffusion
- Stick Slip
- Asperities (kinks)
- Crack Propagation

7. Predicting earthquakes is not yet possible

Seismic Gaps



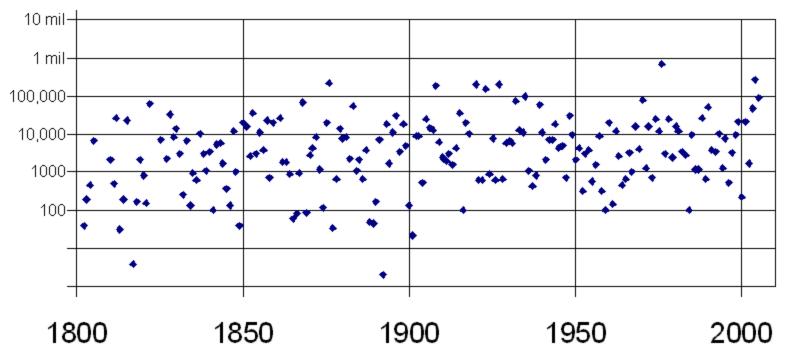
Are Earthquakes Getting More Frequent?



7. Predicting earthquakes is not yet possible

Earthquake Fatalities Since 1800

Fatalities Per Year 1800-2005



7. Predicting earthquakes is not yet possible

Take-Away Points

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